



# Radiometry on Argo Floats: From the Multispectral State-of-the-Art on the Step to Hyperspectral Technology

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Over the past two decades, robotic technology such as Argo floats have revolutionized operational autonomous measurement of the oceans. Recently, Biogeochemical Argo floats (BGC-Argo floats) have measured optical and biogeochemical quantities down to a depth of 2.000 m. Among these parameters, are measurements of the underwater light field from which apparent optical properties (AOPs), such as the diffuse attenuation coefficient for downwelling irradiance  $K_d(\lambda)$ , can be derived. Presently, multispectral observations are available on this platform at three wavelengths (with 10-20 nm bandwidths) in the ultraviolet and visible part of the spectrum plus the Photosynthetically Available Radiation (PAR; integrated radiation between 400 and 700 nm). This article reviews studies dealing with these radiometric observations and presents the current state-of-the-art in Argo radiometry. It focus on the successful portability of radiometers onboard Argo float platforms and covers applications of the obtained data for bio-optical modeling and ocean color remote sensing. Generating already high-quality datasets in the existing configuration, the BGC-Argo program must now investigate the potential to incorporate hyperspectral technology. The possibility to acquire hyperspectral information and the subsequent development of new algorithms that exploit these data will open new opportunities for bio-optical long-term studies of global ocean processes, but also present new challenges to handle and process increased amounts of data.

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# INTRODUCTION

The marine environment is highly complex, and conventional *in situ* measurements carried out, e.g., on research cruises, can only provide a series of snapshots. Therefore, uncovering the real dynamics requires complementation with data of higher spatiotemporal coverage, which could be provided by autonomous measurements (Zielinski et al., 2009; Chai et al., 2020). The Argo network of autonomous floats<sup>1</sup> has revolutionized the capability for operational observations (Roemmich, 2001) to be the main source of CTD measurements (Riser et al., 2016). Biogeochemical Argo floats (BGC-Argo floats) (Biogeochemical-Argo Planning Group, 2016), the extension of the Argo core mission, supply the optical and biogeochemical observations necessary to construct a

<sup>1</sup>https://argo.ucsd.edu

global, full-depth, and multidisciplinary ocean-observing system (IOCCG, 2011; Bittig et al., 2019; Claustre et al., 2020). This type of float accommodates up to six essential biogeochemical and bio-optical variables (**Figure 1**): oxygen (O<sub>2</sub>), pH, nitrate (NO<sub>3</sub>), suspended particles, chlorophyll-a (chl-a), and radiometry.

The basic radiometric measurement is the spectral radiance,  $L(\lambda)$ , representing the radiant flux from a defined direction through a defined area. From radiance, all other integral measures of the light field can be derived, such as irradiance  $E(\lambda)$ ,which is the total radiant flux per unit area of horizontal surface measured in the direction of the incoming light. The downwelling irradiance  $E_d(\lambda)$ , thus the irradiance measured toward the sea surface, is the parameter that is regularly measured by the BGC-Argo floats (Organelli et al., 2016). However, some initiatives have started to look also at the upwelling radiance  $L_u$  (Gerbi et al., 2016; Leymarie et al., 2018; Wojtasiewicz et al., 2018). The propagation of  $E_d(\lambda)$  in the water column can be described approximately as an exponential function of depth (Mobley, 1994):

$$E_{d}(\lambda, z) = E_{d}(\lambda, 0^{-}) e^{\int_{z}^{0} -K_{d}(\lambda, z).dz}$$
(1)

In case a layer-averaged  $K_d$  (from 0 to z) is used, light decline is calculated according to:

$$E_{d}(\lambda, z) = E_{d}(\lambda, 0^{-}) e^{-K_{d}(\lambda, z) \cdot z}$$
(2)

There, z [m] is water depth,  $E_d(\lambda, 0^-)$  [Wm<sup>-2</sup>sr<sup>-1</sup>nm<sup>-1</sup>] the irradiance right below the surface, and  $K_d(\lambda, z)$  [m<sup>-1</sup>] the diffuse attenuation coefficient of downwelling irradiance, which can be deduced from measurements of  $E_d$  in different depths. Both  $E_d$  and  $K_d$  are spectral quantities, indicated by  $\lambda$  [nm] for wavelength.

The slope of this exponential decline, thus  $K_d(\lambda)$ , depends on the absorption (a) and backscattering properties  $(b_{\rm b})$  of the water (Lee et al., 2005), i.e., termed the inherent optical properties (IOPs), which are in turn determined by the water molecules themselves as well as the properties and concentration optically active substances (OAS) within the water column, such as phytoplankton, detrital matter, and colored dissolved organic matter (CDOM; see also Kirk, 2011). From  $E_d(\lambda)$  and  $L_{\mu}(\lambda)$ , other quantities, commonly known as apparent optical properties (AOPs), can be derived. In addition to  $K_d(\lambda)$ , a frequently derived AOP is radiance reflectance,  $r_{rs}(\lambda)$ , (or remote sensing reflectance,  $R_{rs}(\lambda)$ , when measured just above the sea surface). It is defined as the ratio of  $L_u(\lambda)/E_d(\lambda)$ (see Morel and Gentili, 1996). A multiplication with the extraterrestrial solar spectral irradiance, gives another important parameter, the normalized water-leaving radiance,  $L_{wn}(\lambda)$  (see Gordon and Clark, 1981).  $K_d(\lambda)$ ,  $r_{rs}(\lambda)$ , and  $R_{rs}(\lambda)$  are also the most commonly studied parameters of interest in remote sensing applications, notably satellite observation (Austin and Petzold, 1981; Werdell and Bailey, 2005). However, converting a radiometric signal captured by a spaceborne sensor to an enduser product carries uncertainties. High demands for accuracy in absolute calibration of satellite remote sensing data (IOCCG, 2013) and high quality field observations are required for satellite product validation/evaluation (Mueller and Austin, 1992; Zibordi and Voss, 2014). For both optical in situ and remote sensing observation, various bio-optical algorithms have been developed to relate features of the water signal to in-water quantities of interest (e.g., O'Reilly et al., 1998; Maritorena et al., 2002). Improving these retrieval algorithms (existing) and develop enhanced ones that exploit forthcoming technologies remains a priority for ocean optics community.

Ocean optics research has recently undergone a shift from multispectral (with broadband data) to hyperspectral instrumentation, both in situ and remotely (Chang et al., 2004). The higher number of available bands in hyperspectral observations allows a more detailed examination regarding spectral resolution of AOPs and IOPs within the water column. This makes applications requiring higher resolution feasible, e.g., the discrimination of phytoplankton groups, the determination of particle size, and more detailed determination of CDOM spectral properties. Conventional Argo floats already provide sustained high-quality data at unprecedented temporal and spatial resolution. Thus, the addition of hyperspectral observation capabilities will allow Argo platforms to meet demands regarding more sophisticated approaches from the field of optical models and remote sensing. The exploitation of such high-resolution data (in the spatial, temporal, and spectral dimension), measured by a single system that carries also IOP sensors, will certainly open new avenues for bio-optical research in the future. This exciting combination (Argo platform + hyperspectral radiometers) could be the next logical step for Argo floats network (Claustre et al., 2020). With this in mind, this review summarizes the current state-of-the-art of radiometric measurements on Argo floats, and identifies emerging future pathways for the development of hyperspectral Argo radiometry and data exploitation. In our search, we referred to the list of the BGC Argo mission bibliography,<sup>2</sup> from which we selected studies involving radiometric measurements (including papers describing data management approaches and quality-control procedures). We made additional research using Google Scholar and Scopus for verification.

### ARGO-RADIOMETRY: STATE-OF-THE-ART

### Instrumentation

The radiometer mounted on BGC-Argo floats at the time of this review is the multispectral Ocean Color Radiometer (OCR-504, SATLANTIC Inc./Sea-Bird Scientific, United States) (SATLANTIC, 2013), which measures irradiance commonly at four channels. Three channels are usually dedicated to narrow-wavelengths (around 10 or 20 nm bandwidth) from the ultraviolet to the visible range of the spectrum, while the fourth channel measures the Photosynthetically Available Radiation (*PAR*; integrated radiation between 400 and 700 nm). Two different sets of waveband configurations have been documented in the literature. Depending on the funding program, the earlier studies have applicated 412, 490, and 555 nm (Xing et al., 2011, 2012, 2014; Mignot et al., 2014). Later studies have selected

<sup>&</sup>lt;sup>2</sup>https://argo.ucsd.edu/outreach/publications/biogeochemical-argo-bibliography



380 nm instead of 555 nm (Organelli et al., 2016; Organelli et al., 2017a,b; Mignot et al., 2018; Organelli and Claustre, 2019). The fourth channel dedicated to PAR was used in the context of determining overall light available for primary production in the ocean (Lacour et al., 2017, 2019; Mignot et al., 2018; Barbieux et al., 2019; Terzić et al., 2019). However, studies that measured both  $E_d$  and  $L_u$  have used the four wavelengths centered at 412, 443, 488, and 555 nm (Gerbi et al., 2016; Wojtasiewicz et al., 2018), and up to 7 wavelengths plus *PAR* using OCR507 (see Leymarie et al., 2018). **Table 1** summarizes available studies that made use of radiometric data obtained with Argo floats, with the exception of studies that exploit the data only as additional or support information. We also excluded studies that covered primarily data management or quality control protocols.

# Exploitation of the Radiometric Data Obtained From Argo

The mapping of phytoplankton is an area of environmental monitoring that has particularly benefitted from the Argo radiometric observations. Indeed, the *PAR* observations acquired in high sampling frequency provide a comprehensive picture of the vertical distribution of phytoplankton chlorophyll over the different layers in the euphotic zone (surface layer, mixed layer, and deep chlorophyll maximum layer) (Mignot et al., 2014; Lacour et al., 2017, 2019; Barbieux et al., 2019; Ricour et al., 2021). Gaining such dynamic illustration of the light gradient, including the critical depth of the upper euphotic mixed layer, revealed the magnitude and timing of blooms

occurring in otherwise inaccessible locations in different seasons (Randelhoff et al., 2020). However, with respect to phytoplankton diversity, these ecological events are still in demand for a monitoring in terms of pigment groups or functional types (Kutser, 2009), for which multispectral data provide not enough spectral resolution. Readily, this dataset has demonstrated also its utility in improving numerical biogeochemical modeling (Terzić et al., 2019). Other reviewed studies provided that PAR and/or Ed have also successfully used to validate IOP parameterizations (Lazzari et al., 2021). Biogeochemical quantities like chl-a and CDOM were also retrieved from  $K_d(490)$  and  $K_d(412)$ , respectively (Xing et al., 2011, 2012), and other related studies disclosed their main drivers (Organelli and Claustre, 2019). The estimations of such biogeochemical quantities are still in need to be refined and improved better resolved with respect to their relative contribution to the in-water optical properties via their absorption and scattering signatures (Claustre et al., 2020). Acquiring radiometric data over the entire spectrum will permit the construction of continuous spectral diffuse attenuation information that, in turn, enables for depth-dependent IOPs characteristics of the OAS in the water column. Such analysis would be rather difficult with the fewer bands of a multispectral sensor (Chang et al., 2004).

Argo radiometric data plays further a significant role in increasing the amount of data available for satellite matchups in remote areas (Organelli et al., 2017a; Wojtasiewicz et al., 2018), as well as for evaluating the performance of ocean color remote sensing retrievals (Organelli et al., 2017a; Xing et al., 2020). Thus, a further exploitation of the Argo platform capabilities is

Type of Argo float	Optical quantity	Derived AOPs	Area studied	Aim	Author
Not specified	PAR		Black Sea	Vertical distribution of chlorophyll concentration	Ricour et al., 2021
Proice	$E_d$ (380, 412, 490) + PAR		Arctic	Study the phytoplankton growth	Randelhoff et al., 2020
Not specified	<i>E</i> <sub>d</sub> (380, 412, 490) + <i>PAR</i>		Mediterranean Sea	Validation of numerical model of radiative transfer	Lazzari et al., 2021
Not specified	E <sub>d</sub> (380, 412, 490) + PAR	K <sub>d</sub> (PAR); K <sub>d</sub> (490)	World's ocean (Southern Ocean, Subpolar Gyre, Transition zone, Red Sea, Black Sea, Mediterranean Sea, Subtropical Gyre, Arctic Sea, New Caledonia)	Evaluation of ocean color remote sensing algorithms for diffuse attenuation and optical depths	Xing et al., 2020
PROVOR (NKE Marine Electronics Inc., France)	E <sub>d</sub> (380, 412, 490) + PAR	K <sub>d</sub> (380); K <sub>bio</sub> (380); K <sub>star</sub> (380); K <sub>d</sub> (490); K <sub>bio</sub> (490)); K <sub>star</sub> (490)	North Atlantic subtropical gyre	Understand the potential driver of CDOM	Organelli and Claustre, 2019
Not specified	PAR	$K_d$ (PAR)	Mediterranean Sea	Improve the simulation of chl-a dynamics	Terzić et al., 2019
PROVOR (NKE Marine Electronics Inc., France)	PAR	Euphotic depth (Zeu)	North Atlantic subpolar Ocean	Light penetration information	Lacour et al., 2017, 2019
Navis (Sea-Bird Scientific, Inc., United States)	<i>E<sub>d</sub></i> (412, 443, 490, 555) <i>L<sub>u</sub></i> (412, 443, 490, 555)	R <sub>rs</sub> (412, 443, 490, 555)	Indian Ocean subtropical gyres	Validation of remote sensing reflectance	Wojtasiewicz et al., 2018
PROVOR (NKE Marine Electronics Inc., France)	$E_{d} (380, 443, 490, 510, 560, 665) + PAR$ $L_{u} (380, 443, 490, 510, 560, 665)$ $E_{d} (412, 443, 490, 510, 560, 665) + PAR$ $L_{u} (380, 443, 490, 510, 560, 665)$		Mediterranean Sea Southern Ocean	Test the applicability of the new PROVOR float (CTS5) in ocean color validation activities	Leymarie et al., 2018
Not specified	PAR	K <sub>d</sub> (PAR)	North Atlantic subpolar Ocean	Determine the euphotic layer depth	Mignot et al., 2018
PROVOR CTS-4 (NKE Marine Electronics Inc., France)	E <sub>d</sub> (380, 412, 490)+ <i>PAR</i>	К <sub>d</sub> (380); К <sub>d</sub> (490)	Black Sea; Subtropical gyres (Atlantic and Pacific); High latitudes (North Atlantic and southern oceans); Mediterranean Sea	Reexamine the regional variability of relationship between CDOM and phytoplankton light absorption properties referring to existing bio-optical models	Organelli et al., 2017b
APEX (Teledyne Webb Research., United States)	<i>E<sub>d</sub></i> (412, 443, 488, 555) <i>L<sub>u</sub></i> (412, 443, 488, 555)	R <sub>rs</sub> (412, 443, 488, 555)	Mediterranean Sea; Pacific Ocean; Atlantic Ocean	Validation of remote sensing reflectance	Gerbi et al., 2016
PROVOR (NKE Marine Electronics Inc., France)	<i>E<sub>d</sub></i> (412, 490, 555)	K <sub>d</sub> (412); K <sub>bio</sub> (412)); K <sub>bio</sub> (490)	Mediterranean Sea; North Atlantic; North Pacific gyre	Quantification of biogeochemical quantities	Xing et al., 2011, 2012

reasonable to enhance the utilization of satellite data in terms of product retrievals (with less uncertainties). A frequent requested product in this respect are e.g., the phytoplankton accessory pigments in order to characterize phytoplankton dynamics in more detail instead just in terms of biomass (Mueller and Fargion, 2002). Other reviewed papers here have implicated the radiometric Argo data as a metric to get supplementary information about the euphotic zone and optical depth (Doxaran et al., 2014; Mayot et al., 2017), and to resolve events undetectable by satellite observations or classical sampling, for example, variation in optical depth preceding and following a typhoon event (Doxaran et al., 2014).

Other gains sum up in inclusion this dataset in those procedures of fluorometer (accompanied sensor on Argo float) calibration. For instance, measurement of  $E_d$ (490) enables the retrieval of a scaling factor required for the calibration of the chl-a fluorometers (Xing et al., 2011). The reliability of this method was supported by later works (Xing et al., 2012, 2014; Mignot et al., 2014; Ardyna et al., 2019; Kubryakov et al., 2019). Argo *PAR* profiles have been also used to ameliorate the non-photochemical quenching (NPQ) correction of Argo chl-a fluorescence measurements (Xing et al., 2018).

## **Data Availability and Quality Control**

As with all BGC-Argo variables, radiometric datasets are available within 24 h of transmission to the public in the Argo Global Data Assembly Centers (GDACs) along with the corresponding CTD data (Wong et al., 2020). The international Argo Data Management Team (ADMT) supervises the development of data management procedures of Argo data (Argo Data Management Team, 2021). As expected, this review observed a regular protocol in the way of acquiring and processing radiometric data (Schmechtig et al., 2017). In terms of quality control, two levels are generally applied: the "real time mode" that effectuates quick automatic checks (e.g., global range test) to ensure data provision to the user in near real time (Poteau et al., 2019); and the "delayed mode" based on scientific judgement, to account for confounding effects such as biofouling and sensor drift (Schmid et al., 2007), as well as other sources of uncertainty, e.g., the temperature effect on dark counts (Organelli et al., 2017a; Wojtasiewicz et al., 2018). In the earliest stage of the BGC-Argo program, Xing et al. (2011) relied on interpreting the chl-a fluorescence profiles that accompanied the downwelling irradiance measurements to obtain noise-free radiometric profiles. Later, Organelli et al. (2016) developed a quality-control procedure based on statistical and mathematical analyses that targets the major issues related to vertical in-water light field measurements (e.g., dark signal, clouds, wave-focusing, general spike detection; see also Mueller et al., 2003), taking into account the specific circumstances of the Argo platform (e.g., lack of operator control and float recovery precluding post-deployment check or recalibration of the radiometer). As no above-water measurements are available to validate the subsurface radiometric measurement, the comparison of the near-surface measurements with the output of clear-sky optical models (Gregg and Carder, 1990) was suggested to evaluate the obtained measurements (Xing et al., 2011; Organelli et al., 2016).

# **FUTURE PROSPECTS**

Argo radiometric measurements are currently limited to observations measuring, depending on configuration, either irradiance or radiance at up to four channels. Acquisition of radiometric observations takes place in discrete spectral bands, selected to detect major variations in underwater optical properties, allowing these multispectral devices to be effective in deriving the relative concentration of in-water constituents over the water column. Taking the example of phytoplankton, the retrieval of more details regarding species composition (which is oftentimes called for) has been challenging because of the insufficient spectral resolution of the multispectral sensor (Bracher et al., 2017). Hyperspectral radiometers successfully achieved this discrimination (e.g., Isada et al., 2015; Xi et al., 2015), as well as the retrieval of phytoplankton functional types (Wolanin et al., 2016), as they provide measurements at hundreds of contiguous wavebands distributed equally from the ultraviolet to the near-infrared segment of the spectrum with a resolution of the order of 5 nm. This has led to more frequent use of hyperspectral instrumentation in ocean optics research (Chang et al., 2004). Bringing hyperspectral devices into play on robotic platforms such as Argo floats is therefore the next logical step (Claustre et al., 2020). In some, especially remote areas affected by regularly occurring (harmful) algal blooms, monitoring strategies require increased observing capabilities (Sellner et al., 2003). Whilst detection and characterization of the spatial and temporal development of phytoplankton biomass is already feasible using existing multispectral Argo floats, the transformation toward hyperspectral technology will provide additional opportunities of tracking its composition and predict harmful algal bloom development (Shen et al., 2012; Dierssen et al., 2020). Ultimately, hyperspectral data based on Argo float may pave the way toward harmful algal bloom forecasting programs that require nearly real time observations from a global communication network (Jochens et al., 2010).

Hyperspectral data is further advantageous for calibration and validation purposes (McClain and Meister, 2012) of current and future hyperspectral ocean color missions such as the Environmental Mapping and Analysis Program (EnMAP) (Guanter et al., 2015), the Hyperspectral Precursor of the Application Mission (PRISMA) (Loizzo et al., 2018), and the Pre-Aerosol, Clouds, and ocean Ecosystem (PACE) mission (Gregg and Rousseaux, 2017). Indeed, hyperspectral radiometric observations measured by Argo will provide high-resolution data (in the context of spatial, temporal and spectral resolution) that will be more appropriate to match the spectral bands of these advanced sensors (Werdell et al., 2013; Brewin et al., 2016). This would subsequently improve the assessment of derived products and thereby upgrade the validation metrics (Maritorena and Siegel, 2005; Maritorena et al., 2010; Seegers et al., 2018). A particular benefits of such datasets would be that they offer the possibility to develop new algorithms with more freedom in selecting spectral bands for bio-optical studies. Such developments would in turn improve both empirical remote-sensing (Gordon and Morel, 1983; Lee et al., 1998; Stramska et al., 2003) and semi-analytical approaches (Carder et al., 1999), including the evaluation of radiative transfer modeling simulations (Mobley and Sundman, 2008). This would also serve as a contribution to the tuning of existing (heritage) IOP retrieval approaches (Werdell et al., 2018).

Hyperspectral Argo data, on other hand, will have their own scientific value in lieu of just supporting satellite missions. The exploitation of high-resolution data (in the spatial, temporal, and spectral extension), measured by a single system that carries also IOP sensors, would potentially reduce uncertainties (due to temporal or spatial biases) in existing inverse algorithms (Rehm and McCormick, 2011). Thence, the availability of hyperspectral AOPs quantities will enable the derivation of high spectral resolution IOP data, and this will considerably reduce the ambiguities introduced by the relative contribution of different accessory phytoplankton pigments for example, as well as retrieving CDOM spectral optical properties which bring additional biogeochemical information (Gordon et al., 2009). The investigation of such contemporaneous measurements is demonstrated in the work of Rehm and Mobley (2013) where they estimated absorption and backscattering coefficients to an accuracy of  $\pm 0.01 \text{ m}^{-1}$  and  $\pm 0.0005 \text{ m}^{-1}$  respectively, based on hyperspectral vertical radiometric observations and concurrent measurements of the beam attenuation coefficient at 650 nm in optically simple oligotrophic waters. In the more optically complex waters where OAS do not covary with chl-a, Wollschläger et al. (2020) obtained a moderately complex parametrization of the underwater light field based on hyperspectral light measurements. Such parametrizations could be strongly improved by accounting for seasonal and regional variation based on data coming from hyperspectral Argo-floats. Increasing the number of bands will moreover allow testing of various wavelengths combinations based on the spectral diffuse attenuation to better link IOPs features (Brown et al., 2004), as well as to check for consistency of calibration of fluorescence observations (accompanied sensor on Argo float) (Xing et al., 2011).

The deployment of hyperspectral radiometers on unattended platforms like Argo floats applies higher demands on the capabilities of the sensors than conventional in-water radiometry. Some studies which performed hyperspectral light field measurements using drifting devices (Zielinski et al., 2006; Barnard et al., 2018) have already demonstrated stable shortterm field deployment of autonomous hyperspectral radiance and irradiance measurements. However, both studies highlight the importance of technical features like radiometer power consumption, long-term stability and reliability, standardized interfaces and protocols, as well as intelligent data handling. At present, the first studies are performed to evaluate the use of hyperspectral radiometers on Argo floats. The sensor used for these studies is an improved, 1,000 m-depth rated RAMSES radiometer (TRIOS, Rastede, Germany) designed to collect the incoming irradiance in the spectral range of 320-950 nm, with a spectral resolution of 2.2 nm and a spectral accuracy of 0.2 nm. The sensor also features the new G2 Modbus-interface for easier platform integration (Trios, 2021). An important consideration for Argo-users is that the increased information content increases the demand for data transfer and processing

capabilities. Handling huge amounts of data within a timeframe of 24 h as strived by the Argo community will be a demanding task. In case of the mentioned RAMSES radiometer, the data size is approximately six times that of the OCR-504. Together with the moderately higher power consumption (45 mA compared to 25 mA for the OCR-504), this constrains either the measurement frequency or deployment duration of an Argo float equipped with this sensor. Therefore, reducing power consumption of the sensor and the dimension of the data onboard the platform prior to transmission could be important future steps. This process would probably demand embedded computing power in the sensor and the platform itself, which poses challenges on the commercially available instrumentation.

## CONCLUSION

This review captures the current state of light field measurements onboard Argo floats, describes their application in biooptical ocean studies, and points to future directions in the transition from multispectral toward hyperspectral radiometry. This transition would allow an exploitation of light field measurements far beyond to what is covered to date.

Moving forward in operating Argo radiometry, hyperspectral observations would provide more detailed insights through direct hyperspectral investigation of the depth-resolved light distribution in different waters bodies and would be relevant to ocean color products. The development of novel approaches that exploit this information will open new avenues for bio-optical long-term studies of global ocean change and foster its relevance for carbon cycle investigation by addressing phytoplankton not only in terms of biomass, but also with respect to taxonomic composition. Critically, exploiting these advantages requires the development, in parallel, of new algorithms and procedures for quality assessment, data management, and on-board data reduction, all of which are intrinsically related to the technical features of both sensor and platform.

# **AUTHOR CONTRIBUTIONS**

AJ made the draft version. JW and OZ participated on the outline of the manuscript. All authors contributed to manuscript revisions and agreed to the submitted version.

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**Conflict of Interest:** The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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# APPENDIX

APPENDIX A | Notations used in this review.

Symbol or abbreviation	Significance/units	
L	Radiance (W m <sup>-2</sup> sr <sup>-1</sup> nm <sup>-1</sup> )	
Lu	Upwelling radiance (W m <sup>-2</sup> sr <sup>-1</sup> nm <sup>-1</sup> )	
Ed	Downward irradiance (W m <sup>-2</sup> nm <sup>-1</sup> )	
$L_{WD}(\lambda)$	Water-leaving radiance (W m <sup>-2</sup> sr <sup>-1</sup> nm <sup>-1</sup> )	
$r_{rs}(\lambda)$	Under-water surface remote sensing reflectance (sr <sup>-1</sup> )	
$R_{rs}(\lambda)$	Above-water surface remote sensing reflectance (sr <sup>-1</sup> )	
$K_d(\lambda)$	Diffuse attenuation coefficient at wavelength $\lambda$ (m <sup>-1</sup> )	
$\mathcal{K}_{bio}\left(\lambda ight)$	Non-watery diffuse attenuation coefficient at wavelength $\lambda$ $(m^{-1})$	
$K_{star}(\lambda)$	Chlorophyll-specific attenuation coefficient at wavelength $\lambda$ (units of square meters per milligram of chlorophyll)	
PAR	Photosynthetically Available Radiation ( $\mu$ mole photons m <sup>-2</sup> s <sup>-1</sup> )	
$K_d$ (PAR)	Diffuse attenuation coefficient of PAR (m <sup>-1</sup> )	
chl-a	Chlorophyll-a concentration (mg m <sup>-3</sup> )	
CDOM	Colored dissolved organic matter	
Z <sub>eu</sub>	Euphotic depth (m)	
IOP	Inherent optical property	
AOP	Apparent optical property	